



Geochemistry and petrogenesis of the Late Cretaceous Büyükdüz Gabbro (Trabzon, NE Turkey)

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Abstract

Late Cretaceous I-type plutons are widespread in the Eastern Pontides orogenic belt (EPOB), NE Turkey. Although the majority of the plutons in the EPOB are diorite to granite in composition, very few contains gabbro composition. In this study, we report petrographic and whole-rock geochemical compositions of the Büyükdüz Gabbro. The rocks of the Büyükdüz Gabbro has a uniform composition and is gabbro in composition. The rocks mainly consist of plagioclase, pyroxene, hornblende and Fe-Ti oxide minerals. They are I-type, low to medium-K calc-alkaline, and metaluminous characters, and have low SiO₂ content (51-53 wt.%) and moderate Mg number (Mg# = 54-59). Decreasing CaO, MgO, Fe₂O₃, TiO₂, P₂O₅, Ni, Y and Sr with increasing SiO₂ contents are consistent with fractional crystallization of plagioclase, clinopyroxene and hornblende. The studied rocks are enriched in large-ion lithophile elements (LILEs) and light rare-earth elements (LREEs). Chondrite normalized REE patterns of the samples are concave shaped (La_N/Yb_N = 3.8-8.8) and display negative Eu (Eu_N/Eu* = 0.77-0.84) anomalies. All these data with regional geology indicates that the parental magma of the studied rocks derived from an enriched lithospheric mantle source in an arc setting.

Keywords: *Büyükdüz Gabbro, Eastern Pontides (NE Turkey), Late Cretaceous, Geochemistry, Petrogenesis*

1. Introduction

The EPOB in NE Turkey, located in the Alpine-Himalayan orogenic belt, is one of the well-preserved continental magmatic arcs [1,2] where plutonic and volcanic rocks are widely observed [3–9]. The plutonic rocks varying ages from Early Carboniferous to Eocene are located in this region. These plutons were emplaced in four different time periods: Early Carboniferous, Middle-Late Jurassic, Late Cretaceous and Eocene. The compositions of these plutons vary from gabbro to granite, although gabbro compositions are not common.

Numerous geochemical and petrological studies have been conducted on Late Cretaceous plutons of the EPOB [10,11]. However, studies on the Büyükdüz Gabbro are scarce [12]. In this study, we report petrographic and whole-rock geochemical compositions of the Büyükdüz Gabbro to determine their magma parental sources.

2. Material and Method

The whole-rock geochemical analysis of the rock samples, which were compiled from the study area, were carried out in the ACME Analytical Chemical

Laboratory (Vancouver, Canada). Major and trace elements were analyzed by ICP-AES method and rare earth elements (REEs) were analyzed by ICP-MS method. Details of analysis technique are given in [11].

3. Regional and local geology

The Eastern Pontides is commonly subdivided into southern and northern zones [2,13]. The study area is located in the northern zone of the EPOB (Fig. 1).

The pre-Late Cretaceous units in the EPOB consist mainly of Early Carboniferous metamorphic rocks [14], Late to Early Carboniferous plutonic rocks [15–22], Early and Middle Jurassic volcano-sedimentary rocks [23,24], Mid to Late Jurassic plutonic rocks [25–28] and Late Jurassic to Early Cretaceous carbonates [29]. Late Cretaceous units, including the studied Büyükdüz Gabbro, consist of plutonic, volcanic and sedimentary rocks [30–37]. Cenozoic units consist of plutonic-volcanic (adakitic and non-adakitic) and sedimentary rocks [38,39,48,49,40–47]. Quaternary alluvium forms the youngest units in the region.

The oldest units in the studied area comprise Liassic aged Hamurkesen Formation consisting of mainly basalt, andesite and pyroclastic rocks. Late Cretaceous aged Çatak Formation consisting of andesite, basalt and pyroclastic rocks lie unconformably on these units. All these units were cut by Late Cretaceous aged Camiboğazı Pluton and Büyükdüz Gabbro.

Büyükdüz gabbro is approximately ellipse-shaped and the long axis extends in the northwest-southeast direction (Figure 1). It is dark colored in the field and is rich in mafic minerals.

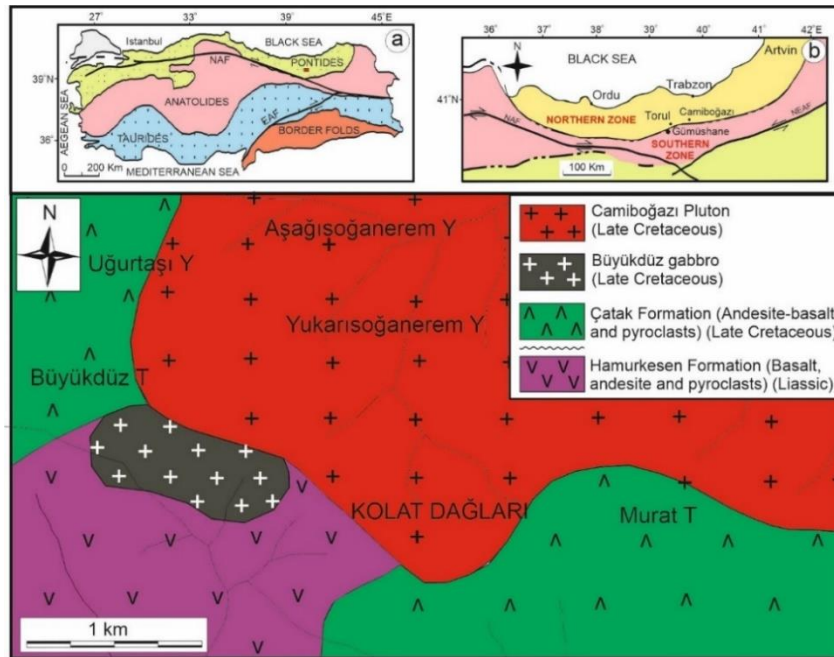


Figure 1. Geological map of the study area (modified after [11])

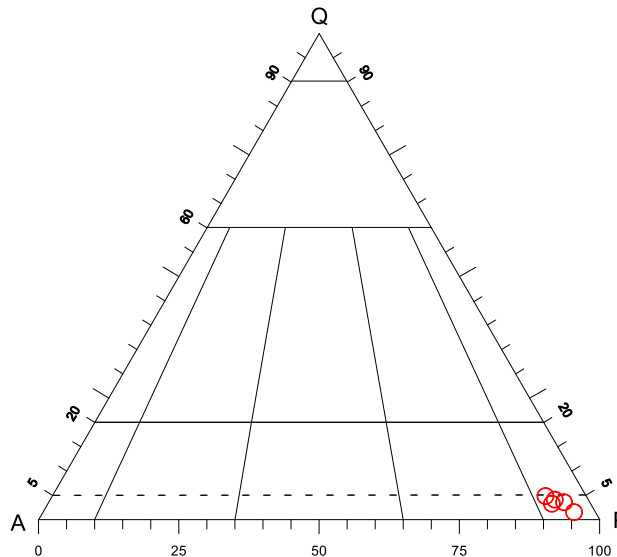


Figure 2. Classification based on modal compositions of the Büyükdüz gabbroic rocks [50].

4. Results

4.1. Petrography

The Büyükdüz gabbroic rocks are gabbro in compositions in the modal QAP diagram (Figure 2) [50].

The studied gabbroic rocks are generally holocrystalline, medium to coarse grained. All samples from the studied rocks consist of equigranular plagioclase (modal 68 to 71%), clinopyroxene (6 to 9%) and hornblende (10 to 12%) as major mineral phases. Apatite, zircon and Fe-Ti oxides form accessory minerals. Secondary phases comprise sericite, calcite, chlorite and clay minerals (Table 1). Plagioclases form euhedral and subhedral lath-shaped phenocrysts. Clinopyroxenes are found as subhedral to anhedral, weakly-zoned grains and also in aggregates. Hornblends are in the form of euhedral

prismatic crystals and contain small apatite and opaque minerals as inclusions. They have well developed cleavages. Subhedral to anhedral apatites are relatively abundant accessory minerals. Zircons are seen as small euhedral prismatic crystals. Opaque minerals are found as subhedral and euhedral in both small and large crystals.

4.2. Whole-rock Geochemistry

4.2.1. Major and trace elements

Considering the results of the geochemical analysis, the SiO₂ values of the rock samples from the Büyükdüz Gabbro vary between 51 and 53 wt.%, while their MgO content varies between 4.9 and 6.2 wt.% (Table 1). The K₂O/Na₂O ratios are less than 1 and are between 0.2 and 0.5. A/CNK values are between 0.8 and 0.9, and the magnesium numbers [Mg# = 100 * (MgO / (MgO + Fe₂O₃^T))] are between 54 and 59 (Table 1).

Table 1. The whole-rock major, trace and REE element analysis of the studied samples

	Min	Max	Avg			Min	Max	Avg
SiO ₂	50.96	52.94	51.68		La	9.20	21.90	15.58
TiO ₂	0.42	0.54	0.49		Ce	19.50	50.60	31.28
Al ₂ O ₃	16.93	17.96	17.40		Pr	2.28	5.75	3.82
Fe ₂ O ₃ ^T	8.42	9.45	9.01		Nd	10.00	23.80	16.68
MnO	0.16	0.19	0.18		Sm	2.35	4.20	3.42
MgO	4.86	6.17	5.41		Eu	0.67	0.99	0.86
CaO	7.53	9.71	8.81		Gd	2.51	3.28	2.96
Na ₂ O	2.12	2.46	2.29		Tb	0.42	0.55	0.49
K ₂ O	0.34	1.34	0.94		Dy	2.66	2.92	2.77
P ₂ O ₅	0.05	0.09	0.07		Ho	0.55	0.67	0.60
LOI	3.40	3.70	3.52		Er	1.52	1.75	1.65
Total	99.73	99.85	99.78		Tm	0.21	0.27	0.24
Ni	9.20	15.80	11.74		Yb	1.63	1.84	1.72
V	149.00	213.00	190.40		Lu	0.24	0.27	0.25
Cu	26.90	78.60	55.62					
Pb	7.80	19.60	14.28		K ₂ O/Na ₂ O	0.16	0.54	0.40
Zn	36.00	59.00	53.00		Mg #	54.40	59.26	56.87
W	0.50	0.70	0.60		A /CNK	0.79	0.94	0.84
Rb	6.60	35.80	20.52		Sr/Y	17.06	25.84	21.07
Ba	249.00	470.00	355.20		Nb/Ta	9.33	37.00	17.97
Sr	244.00	387.60	322.60		Y/Nb	3.67	6.52	4.82
Ta	0.10	0.30	0.22		Nb/La	0.11	0.39	0.24
Nb	2.50	3.90	3.30		Ce/Pb	1.24	6.49	2.63
Hf	1.30	2.20	1.84		Zr/Sm	18.45	26.22	23.06
Zr	57.40	106.20	78.38		(La/Lu) _n	3.97	8.72	6.32
Y	14.30	16.30	15.30		(La/Sm) _n	2.46	3.28	2.81
Th	2.70	3.00	2.84		(Gd/Lu) _n	1.29	1.62	1.45
U	0.70	0.90	0.78		(La/Yb) _n	3.81	8.81	6.09
Ga	13.50	15.80	14.16		(Eu/Eu*) _n	0.77	0.84	0.81
Mg# (mg-number) = molar 100xMgO/(MgO+Fe ₂ O ₃ ^T). LOI = loss on ignition, Eu* = (Sm+Gd) _N /2								

In the SiO_2 versus $(\text{Na}_2\text{O}+\text{K}_2\text{O})$ diagram (Figure 3a), the majority of the samples plot on the gabbro area and a little gabroic diorite area (Figure 3a). All of the

samples are subalkalen in characters and have low to medium-K content in the K_2O versus SiO_2 diagram (Figure 3b).

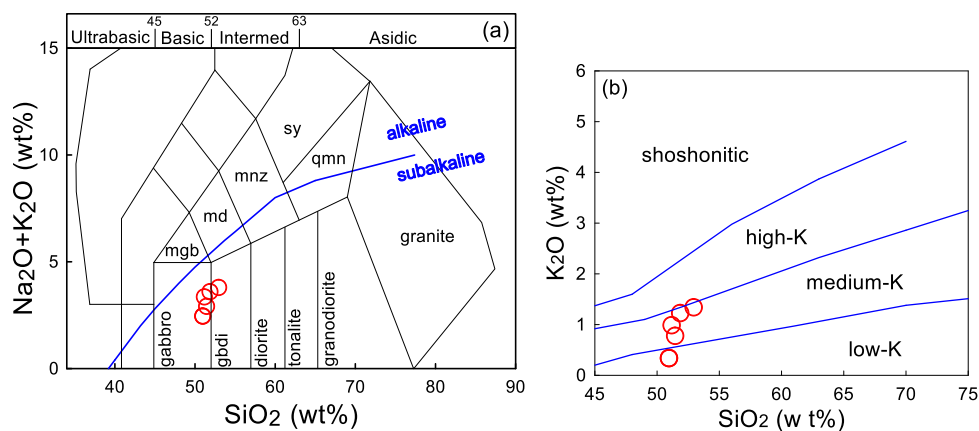
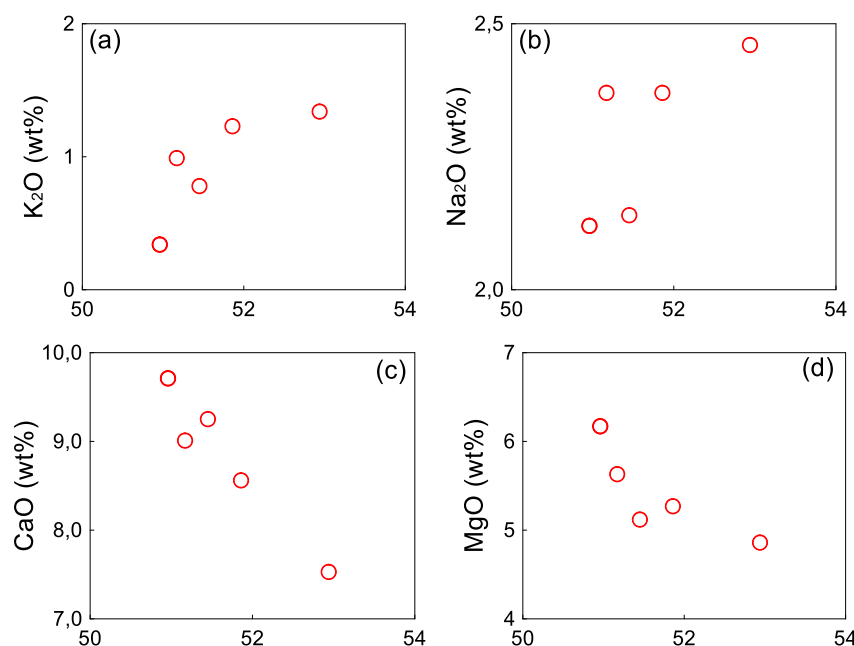


Figure 3. (a) $(\text{Na}_2\text{O}+\text{K}_2\text{O})$ versus SiO_2 diagram [51], (b) K_2O versus SiO_2 diagram [52] for the samples from the Büyükdüz Gabbro

Major and trace element exchange diagrams (Harker diagrams) are given in Figure 4 and 5. In the major elements versus SiO_2 diagrams, Fe_2O_3 , MgO , CaO , P_2O_5 and TiO_2 decrease, and K_2O , Na_2O and Al_2O_3 increase with increasing SiO_2 values (Figure 4). In the

trace element versus SiO_2 diagrams, Ni, Sr and Y show a negative correlation with the increasing SiO_2 , while Ba, Rb, Zr, La, Pb and Nb show a positive correlation (Figure 5).



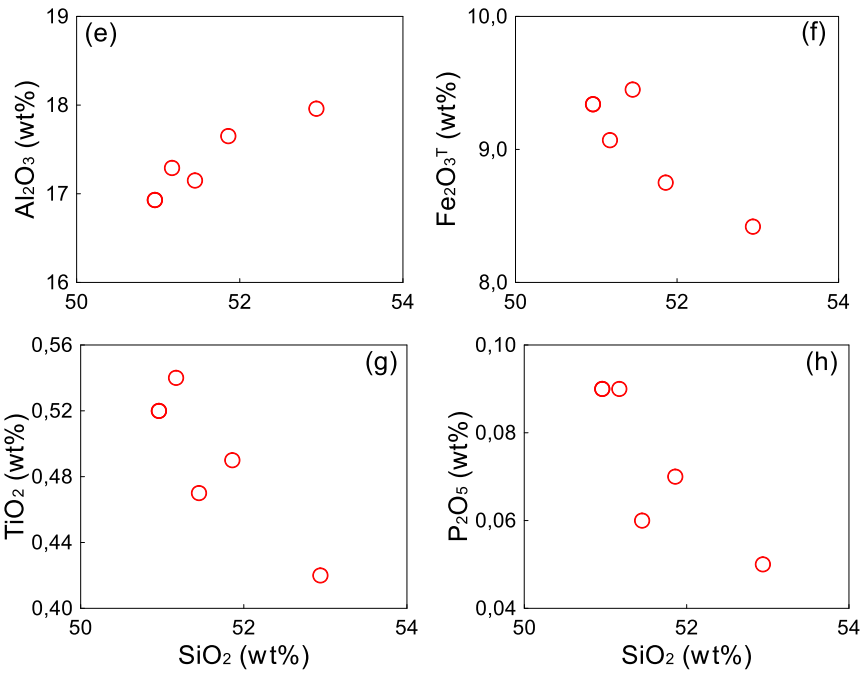
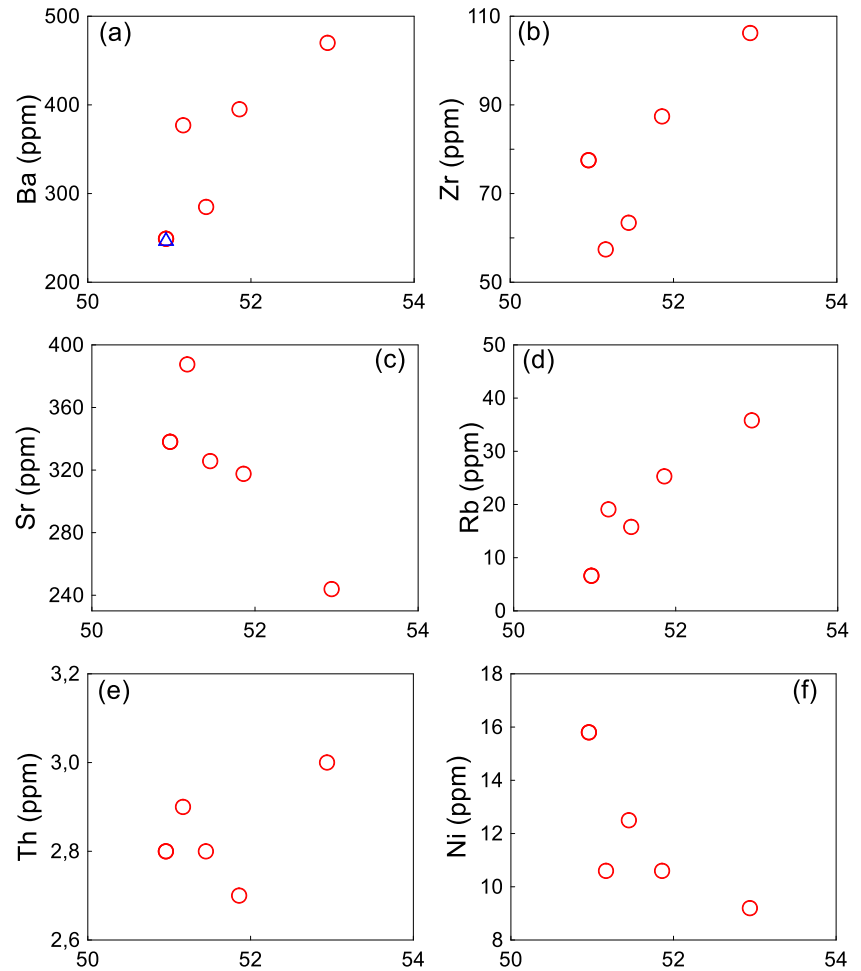


Figure 4. SiO₂ versus major element variation diagrams of the rocks from the Büyükdüz Gabbro



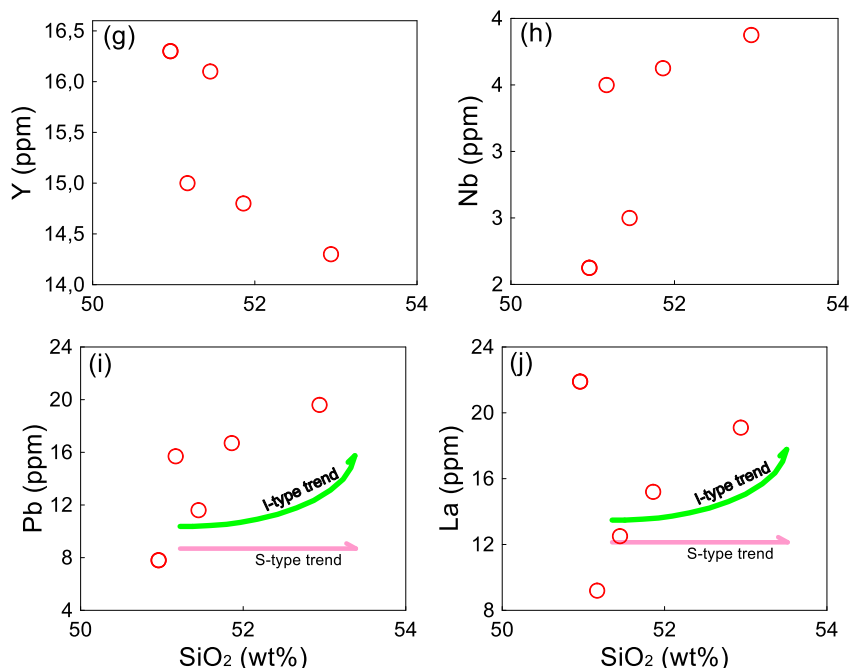


Figure 5. SiO_2 versus trace element variation diagrams of the rocks from the Büyükdüz Gabbro

In the trace element variation diagram normalized to the primitive mantle [53], the samples are characterized by enrichment in LILEs and depletion in high-field strength elements (HFSEs) (Figure 6a). The negative Nb, Ta, P, Ti and positive Pb anomalies are seen in the samples (Figure 6a).

According to chondrite-normalized (normalization

values [54]) rare earth element diagrams (Figure 6b), the studied samples show enrichment in LREEs with $(\text{La}/\text{Sm})_N$ values of 2.46-3.28, and $(\text{La}/\text{Yb})_N$ values of 3.81-8.81. The $(\text{La}/\text{Lu})_N$ values of the samples are between 3.97 and 8.72, and $(\text{Gd}/\text{Lu})_N$ values are between 1.29 and 1.62 (Figure 5b). Eu_N values of the samples are less than 1, and they show negative Eu_N (Eu/Eu^*) anomalies of 0.77-0.84 (Figure 6b).

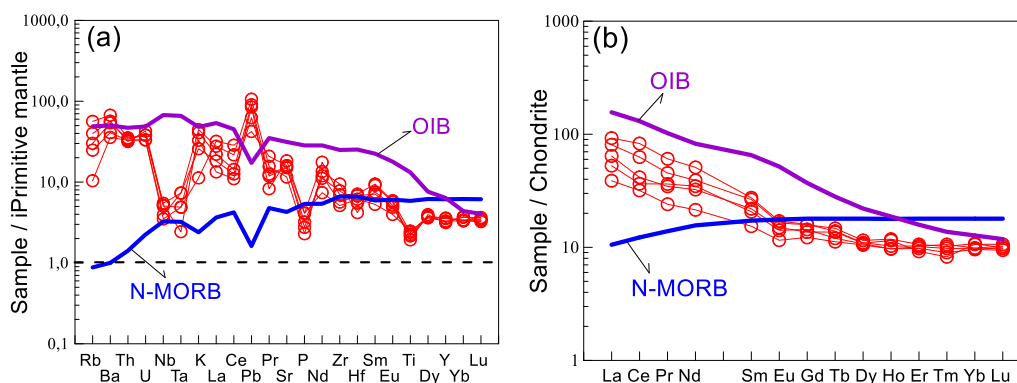


Figure 6. (a) Primitive mantle-normalized [53] trace element patterns, (b) Chondrite-normalized [54] rare earth element patterns for the rock samples from the Büyükdüz Gabbro

4.2.2. Tectonic setting

In the SiO_2 versus ASI (molar $\text{Al}_2\text{O}_3/(\text{CaO}+\text{Na}_2\text{O}+\text{K}_2\text{O})$) diagram, the samples show metaluminous in characters and plot on the I-type field

in the Nb versus $10000\text{GA}/\text{Al}$ diagram (Figure 7 a and b).

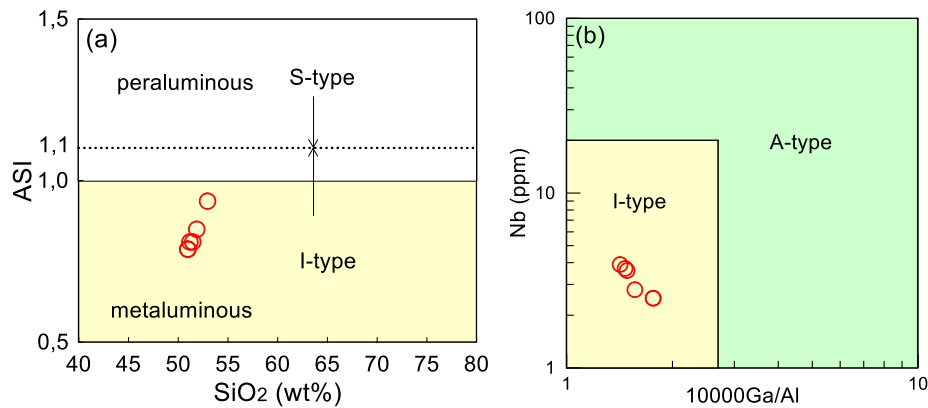


Figure 7. (a) SiO₂ versus ASI diagram (fields: [55,56], (b) Nb versus 10000Ga/La diagram [57] of the samples from the Büyükdüz Gabbro

The Sr/Y ratios of the studied samples are between 17.1 and 25.8, and plot on the volcanic arc field in the Sr/Y versus Y diagram (Figure 8a). In the Rb/Zr

versus Y diagram (Figure 8b), the samples have in the normal arc maturity.

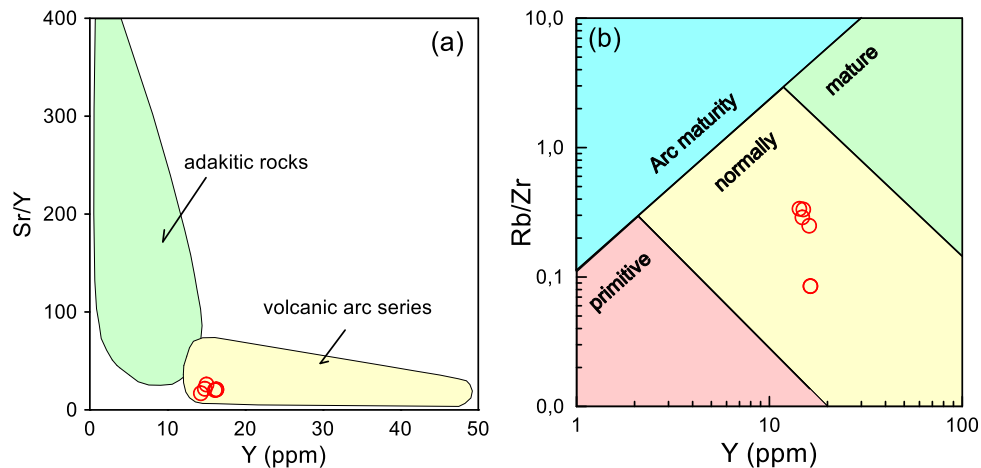
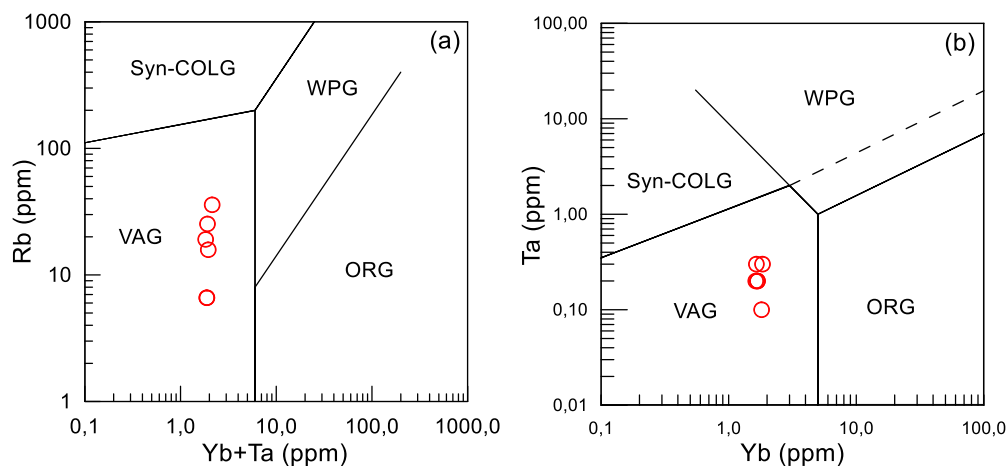


Figure 8. (a) Sr/Y versus Y diagram [58], (b) Y versus Rb/Zr diagram [59] of the samples from the Büyükdüz Gabbro

In the Rb versus (Yb+Ta) diagram [57], the samples fall in the volcanic arc granitoid (VAG) field (Figure 9a). Similarly, they are located in the VAG field in the

Ta versus Yb, Nd versus Y and Ta/Yb versus Ta+Yb diagrams (Figure 9b-d).



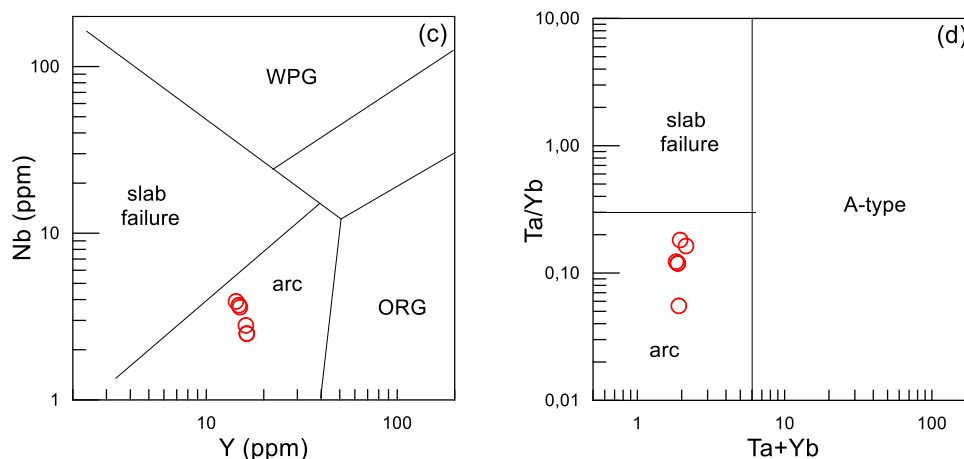


Figure 9. (a) Rb versus (Yb+Ta) diagram, (b) Ta versus Yb diagram [57], (c) Nb versus Y diagram, (d) Ta/Yb versus Ta+Yb diagram of the samples from the Büyükdüz Gabbro

5. Discussion and conclusions

Most of the whole-rock major and trace element variation trends with SiO_2 for the studied samples most likely reflect fractional crystallization. The decreases in MgO , Fe_2O_3 , CaO , P_2O_5 , TiO_2 and Sr with increasing SiO_2 point to a role for the fractionation of plagioclase, clinopyroxene, amphibole and Fe-Ti oxide (Figures 4 and 5).

The Ba/Sr ratios of the studied samples exhibits negative correlation with increasing Sr contents, which can be explained by plagioclase fractionation (Figure 10a). In the CaO versus Y plot [60], all samples has an J-type trend, suggesting clinopyroxene and hornblende fractionation (Figure 10b).

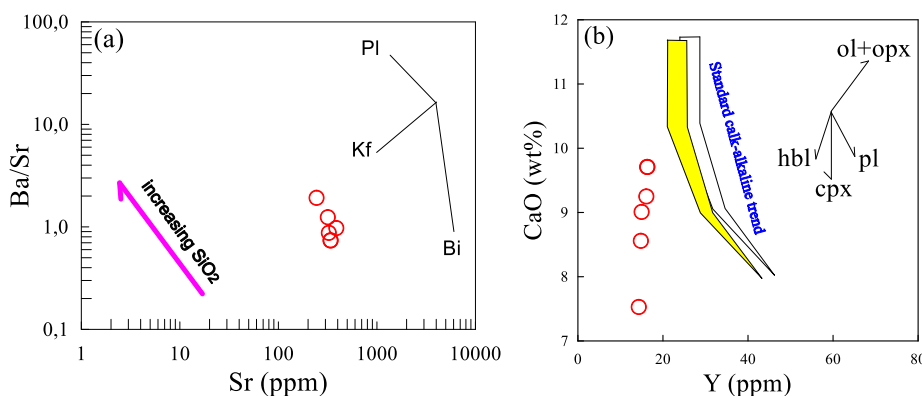


Figure 10. (a) Ba/Sr versus Sr diagram; (b) CaO versus Y diagram [60] from the studied rocks rocks.

The samples of the studied gabbroic rocks are also characterized by negative Nb, Ta, Ti and P and positive Rb, K, Th, and Pb anomalies as well as enrichment in LILEs and LREEs (Figures 6 and 7). These features are commonly found in crustal rocks. Experimental data also show that the rocks formed by partial melts of the lower and middle crust should have Mg numbers ($\text{Mg\#}$) less than 41 in the range of $\text{SiO}_2 \leq 51-65$ [61,62]. The studied samples have low SiO_2 content of 51-53 wt.% and high $\text{Mg\#}$ (53-56). If the source magma was derived directly from partial melting of mafic lower crust, then the primitive samples from the studied rocks should have low $\text{Mg\#}$ [61] in the range of $\text{SiO}_2 \leq 51-65$. However, such is not the case, and the Mg numbers of the studied mafic rocks are clearly higher than those from experimental

crustal melts (<41), as shown in Figure 11a. For this reason, the studied samples may be derived from mantle lithologies, rather than from mafic-felsic crustal lithologies.

Some authors [63,64] suggested that, high Nb/La ratios ($\sim >1$) demonstrate an OIB like asthenospheric mantle source for basaltic magmas, and lower ratios ($\sim <0.5$) display a lithospheric mantle source. The Nb/La (0.1–0.4) ratios of the samples indicate a lithospheric mantle source (Table 1), as shown in Figure 11b. In the Nb/Th versus Nb diagram, all samples plot in the arc fields (Figure 11c). The Th/Yb versus Ta/Yb [65] (Figure 11d) diagram, the studied rock samples form a trend sub-parallel to the mantle array but shifted to higher Th/Yb ratios. This indicates

melt derivation from a source, which had been previously enriched (or metasomatized) by fluids derived from an subduction processes.

All geochemical data has shown that the rocks of the Büyükdüz Gabbro represent a period of (immature)

arc development and probably formed during an early stage of convergent plate boundary tectonics. The magmatism is related to subduction of the Neo-Tethyan Ocean beneath the Eurasian plate along the IAE suture zone during the Late Cretaceous.

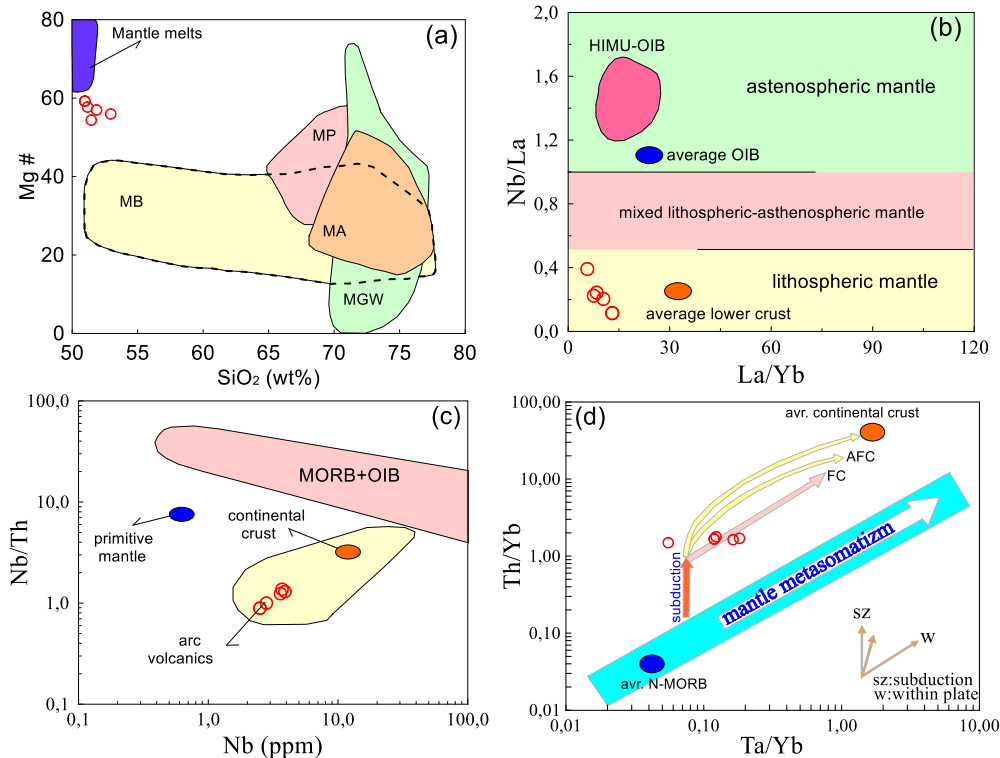


Figure 11. (a) Mg# versus SiO₂, (b) Nb/La versus La/Yb plots, (c) Nb/Th versus Nb, (d) Th/Yb versus Ta/Yb diagrams from the studied rocks. Compositions of continental crust, mid-ocean ridge basalts (MORB), ocean-island basalts (OIB) and arc volcanics after [66]; Average OIB is after [67] and average lower crust is after [68]. Dashed lines separating fields of the asthenospheric, lithospheric and mixed mantle are plotted based on data given in [64].

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